ject, prepared by Elster and Geitel, will be found in Part II ington often hide the summit of the Washington Monument, of the Report of the International Meteorological Congress at Chicago, 1893.

## HURRICANES IN THE WEST INDIES.

From Jos. Ridgway, Jr., St. Thomas, W. I., the Weather Bureau has received the following note, dated October 29:

Thirty years ago we experienced here, on this date, the most severe hurricane ever known in these latitudes. Apparently there is now no further fear of hurricanes for the season 1897, but I may mention as a peculiar feature that on different occasions since the 15th instant the barometer has been lower than at any time during the period 15th July to 15th October.

## ALTITUDES OF CLOUDS.

The well-known voluntary observer, Mr. Barry C. Hawkins, of Horse Cove station, near Highlands, Macon Co., N. C., requests information with regard to the best methods of determining the altitudes of clouds, applicable to the case where the observer is on a mountain or plateau so high that he is sometimes within the cloud. Under date of November 7 he writes:

In mountain regions the altitudes of clouds are not nearly as great as near the sea level, the relation of the observer to the clouds being entirely changed. For instance, the altitude of a cloud near sea level might be 3,000 feet, the same kind of cloud when passing over a mountain range or plateau 3,000 feet high would touch the earth. Such is the actual case in the Blue Ridge Mountains. In storms the clouds frequently sweep the ground, and at other times the height varies from 0 to 500, 800, or 1,000 feet. When touching the earth the cloud appears like fog and might be called fog, but really the case is entirely different, for the origin is like that of all clouds, and the velocity is often great, being equal to that of the surface winds. When the cloud is dense one can only see a few rods. How can we draw a sharp distinction between such clouds and fogs? It is to be noted that cirrus, cirrostratus, and rarely cumulus are so low in altitude. The clouds of thun-In mountain regions the altitudes of clouds are not nearly as great as stratus, and rarely cumulus are so low in altitude. The clouds of thun-derstorms, which are well developed and not connected with any general storm, rarely touch the mountain, peaks of 4,500 feet or more being free from them. But after thunderstorms small clouds, sometimes not free from them. But after thunderstorms small clouds, sometimes not more than 10 feet in diameter, will form perhaps only a few feet above the surface and float in various directions with various velocities, but as the atmosphere clears they disappear. I have sometimes used the following method for obtaining the height of general cloud sheets. Suppose an observer is at an altitude of 2,000 feet and sees a cloud sheet touch the summit of a mountain known to be 3,000 feet high. The altitude of the cloud above will equal 3,000—2,000 =1,000 feet. Ought such altitudes to be counted as above the observer or as above the sea level?

The altitudes of clouds, as published in connection with the international charts of cloud types are given as above sea level, and this is always to be understood when speaking of clouds unless specifically stated to the contrary. Measurements of clouds are made from elevated points, and in publishing such results, both the measured and the sea level altitudes should be given.

The methods of observation and measurement are so varied that students of the subject should refer to the chapters on this subject in the Editor's Meteorological Apparatus and MONTHLY WEATHER REVIEW for April, 1897. The method described by Mr. Hawkins is perfectly proper, and in fact is the one first used by meteorologists in the beginning of the study of this subject; it is quite accurate when there are many clouds whose lower sides are all on the same level, so that an observer by ascending or descending a mountain slope may determine to within a very few yards whether he is on the right level or not. Of course, the actual result depends that becomes warm by compression, is in either case being equally upon the accuracy with which one knows the elevations of his position on the mountain side. This latter con- ders it latent), therefore, the rain, the hail, and the air are sideration is perhaps the most important one when an ob- each cooled in proportion to the quantity of heat that they server looks upward from below and tries to ascertain at what give up to the process of evaporation; of course the air in level the base of the cloud intersects the mountain side. the region at a distance from the falling hail is not cooled.

550 feet above the ground or 600 feet above sea level, but owing to the absence of a sufficient number of well-defined marks on the side of the Monument, as well as to the haziness of the cloud, it is rarely possible to determine the altitude of the base of the cloud to within 50 feet. In the case of observers on the broad plains of the Mississippi watershed it is important that they give the heights of clouds above sea level as well as above ground, since, in general, the types and motions of the clouds depend equally upon the pressure, the temperature, and the moisture of the atmosphere.

## ORIGIN OF THE DESCENDING GUSTS OF WIND.

In the Monthly Weather Review for August, page 351, the Editor has ventured a few words in connection with an extract from a letter of Mr. Charles A. Love, of Aurora, Ill., who now writes that perhaps the wording of his letter has been misunderstood by the Editor, and that he himself had no idea of suggesting that hail causes the wind, but, on the contrary, that cold, dry wind causes the hail. He intended mainly to raise the question as to-

Whether it is possible for a stratum of cold, dry air to get between an upper and lower rain cloud and freeze the rain which is falling from the upper cloud. Is it possible for a stratum of cold air, of relatively greater density, to overlap the warmer air at the surface and cause a downrush and at the same time freeze rain into hail, if there should be any rain in the path of the descending cold air? It is often said that a hailstorm in a certain locality is the cause of the cold, although the hailstorm is of limited area and the change to cold weather covers a wide area. The cold air causes the hail and not the hail the cold air.

The Editor owes Mr. Love an apology for reading too much meteorology into his former letter. The fact is, as we understand it, that when a great mass of cool, dry, and, therefore, denser air rolls southeastward over the continent, displacing warmer, moister air, the process does produce thundershowers and hailstorms, and still oftener raw, cold gusts without much rain or hail. It is, therefore, not proper to say that hailstorms or thunderstorms produced this spell of cool weather, since it was the cold air that caused them, and to this extent our correspondent is quite correct.

Denser air can not lie quietly above lighter air near the earth's surface for any length of time, but if the former is in rapid motion, as is oftentimes the case, bounding along over the earth's surface like the ricochet of a cannon ball, it must necessarily descend in gusts here and there precisely as observed by Mr. Love. At other times the word ricochet does not correctly express the movement, since sometimes the cold wind blows against an obstacle or blows from hilltop to hilltop over a valley before it has time to descend. The descent is usually slow as compared with the horizontal movement.

But there is another equally frequent case in which it is not the cold air that forms the hail, but the hail and rain that forms the cold gust. This latter case is due to two very different modes of action. The first, or simplest of compre-Methods, Washington, 1887, as also to the articles in the hension, is a simple mechanical process; the descending raindrops or hailstones, by the momentum of their fall and the viscosity or internal friction of the air, drive some air before them, so that beneath the column of falling rain the air is pushed outward in all directions, especially southeastward, in the line of least resistance. The second is a thermodynamic action; the falling rain or hail, no matter whether it is entering warmer air or whether it is driving in front of it air evaporated, and as this evaporation consumes heat (i. e., ren-The low clouds that sometimes hang over the city of Wash-The heavy, cold air thus formed descends with the rain and,